

Geochemical Evaluation of Thermal Maturity of the Cretaceous Sediments in the Calabar Flank based on the Distribution of the Polynuclear Aromatic Hydrocarbons

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Abstract: *The thermal maturity parameter used in the evaluation of the hydrocarbon generative potential of the sedimentary organic matter is based largely on the extent of molecular transformation of the biological markers due to geothermal stress which is a function of depth or burial history. In particular the transformations leading to the formation of Polynuclear Aromatics usually occur at the end of diagenesis up to the onset of catagenesis. . An in-depth study of the transformation leading to the formation of polynuclear aromatic steroidal hydrocarbons revealed a step by step thermally mediated aromatization of the non-aromatic biogenic species during diagenetic/catagenetic evolution of organic matter in an anoxic environment. Since the formation of PAHs is a heat mediated process which increases with depth, then its concentration or the amount per unit mass of the rock sample can be used to estimate the thermal maturity of organic matter in the sediment. In light of the above, this work has evaluated the distribution profile of the PAHs in the Cretaceous sediments of the Calabar Flank using Gas Chromatograph interfaced to a Mass Spectrometer (GC-MS), an analytical tool combination widely used in the study of petroleum geochemistry. The results of the study based on the most dominant polynuclear aromatic species namely, anthracene and phenanthrene revealed that Nkporo Shale (NS) samples from various locations in the Basin had a average PAH of 0.05mg/Kg (anthracene) and 0.06mg/Kg (phenanthrene). While the PAH value of Ekenkpon Shale (ES) samples was on the average 0.01mg/Kg and 0.02mg/Kg for anthracene and phenanthrene respectively. New Netim Marl (NNM) however gave the PAH value of only 0.001mg/Kg for phenanthrene and nil for anthracene. On the other hand Awi Sandstone (AS), Mfamosing Limestone (ML) and Nkporo Mudstone (NM) all fell within the non-detectable range. Based on the foregoing, PAH values can be used to evaluate the maturation level of organic matter in the sediments.*

Keywords: *Cretaceous Sediments, Polynuclear Aromatic Hydrocarbons, Thermal Maturity, Distribution Profile*

I. Introduction

The name Calabar Flank was first coined by [1] for the part of the continental margin of Nigeria characterized by block faults with NW-SE trending hoist and graben structures such as Ituk high and the Ikang trough. Thus the Calabar Flank with a total thickness of over 3,500m is related to the South Atlantic Cretaceous marginal basins which are characterized by similar horst structures [2]. The exposed Cretaceous sediments of the basin which is the major focus of this work appear rather unique in that the entire Upper Cretaceous sequence is exposed within a narrow stripe measuring about 8Km (Fig 1b). A great deal of work has been carried out in this marginal sedimentary basin since the last three to four decades and these include works on biostratigraphy, depositional environment and diagenesis of Albian carbonate by [3]; the sedimentology of the Cretaceous sediments by [4]; petroleum source rock assessment of Nkporo shale [5], as well as Upper Cretaceous assemblage from Nkporo Shale by [6] to mention but a few. The present work is aimed examining the distribution of Polynuclear Aromatic Hydrocarbons (PAHs) and the use of PAH concentration as a parameter to evaluate thermal maturity and by extension the hydrocarbon generative potentials of the sediments in the marginal basin. This work is driven by the need to expand the frontier of search for hydrocarbon in the area because other geological provinces such as marginal basins of Equatorial Guinea, Angola and Gabon all in Africa, that show striking similarities in terms of stratigraphy and geological setting with the Calabar Flank are prolific oil fields. But despite long years of search no commercial accumulation of hydrocarbons has so far been discovered in the Calabar Flank. It is hoped that the data obtained in this study will further boost the research efforts in studying the sedimentary geochemistry of the Calabar Flank with the ultimate goal of ascertaining its hydrocarbon generative potentials, both in short and long terms.

1.1 Geological Setting and Stratigraphy

The geological expression known as the Calabar Flank was first described by [1] and later by [7] as an epirogenic basin situated in the easternmost part of the sedimentary basin in southern Nigeria. The Calabar Flank is part of the continental margin of Nigeria dominated by block faults with NW-SE trending horst and graben such as Ituk and the Ikank trough. The basin which lies on Lat. 5°00'N and between Long. 8°00' and 8°30'E (Fig. 1a) contains about four thousand metres of Albian to Maastrichtian sediments in outcrop section [8] and is bounded in the north by the basement rock of Albian Oban Massif; in the south by the Niger Delta hinge line; in the east by the Cameroun volcanic line and in the west by Ikpe platform. [9] suggested a rift origin of the Calabar Flank while [10] sees the basin formation as a response to a slight adjustment in the shape of Africa during its crustal plate separation from South America coupled with the eustatic changes in the sea level within the South Atlantic. Hence its description by [11] as a distinct geologic province with peculiar tectonic style and attendant stratigraphic evolution. The Cretaceous sediment of the Calabar Flank is made up of the Upper and Lower segments. The Upper sediments range from dark gray to bluish shale with mudstone intercalation (Nkporo shale) to thick black shales with mudstone, sandstone and occasional limestone intercalations. While the sediments in the Lower segment are largely flaggy (fissile) limestone with admixture of ferruginized and conglomeratic sequence of Awi Formation. The stratigraphic succession in the Calabar Flank is mostly of Cretaceous age comprising a basal Neocomian-Aptian syn-rift fluvial sandstone, the Awi formation and the marine syn-rift Odukpani group of Albian and Late Cretaceous age. The Odukpani Group [12] comprises the Middle Albian Mfamosing limestone, the Late Albian Ekenkpon shale and the Conician New Netim Marl.

Sedimentation in this marginal basin began with the initial deposition of fluvio-deltaic cross-bedded sands of Awi formation [13] in early Cretaceous times (probably Aptian). This lithologic unit overlies the Precambrian basement complex of Oban Massif. The Awi formation (non-fossiliferous arkosic sandstone), is in turn overlain by a relatively stable carbonate platform, namely Mfamosing Formation, a limestone formation that is partly exposed along the pre-cambrian basal complex of Oban Massif as a result of diatropism (probably an epirogenic uplift) associated with the Neogene tectonics along the Cameroun volcanic line [2]. Thus the Limestone deposits emerged as a result of the first marine incursion in the Middle Albian after which the faulted blocks allow widespread deposition sequence of shales with very minor calcareous (marl) intercalations alternated by the deposition of thin dark shale beds of dominantly shale lithology with occasional mudstone [8]. The lithologic unit immediately overlying this carbonate platform is a thick black and highly fissile shales with minor but frequent intercalation of marls and calcareous mudstone known as Ekenkpon shale formation, believed to be a product of two transgressive cycles that took place during the late Albian-Cenomanian and early Turonian. Ekenkpon (brownish grey/dark grey flaggy shales with mudstone intercalations) formation underlies the New Netim Marl (marlstone with occasional shale intercalations), which together form part of the Odukpani Group, a Late Cretaceous post-rift marine unit [2]. The Awi Formation and the Odukpani Group were deposited before the Santonian compressional tectonic phase [14]; while the Campanian-Maastrichtian Nkporo Shale (bluish black carbonaceous shale with mudstone intercalations), the youngest lithologic unit of the Cretaceous sediments in the stratigraphic succession of the Calabar Flank overlies the New Netim Marl. The sedimentary successions in the Calabar Flank by various workers are presented in (Tables 1-2), with Awi Formation as the oldest at the base and Nkporo shale whose outcrop section remains the youngest lithologic unit of the Cretaceous is at the top [6].

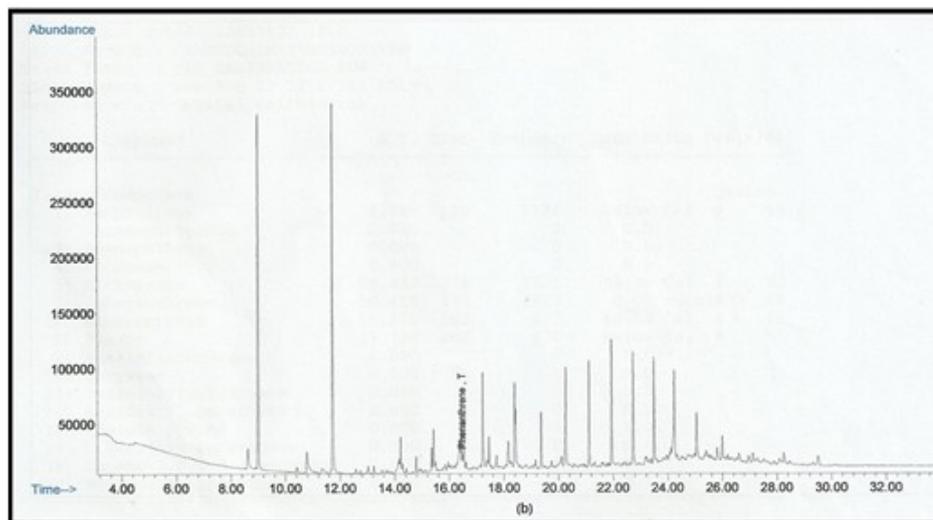
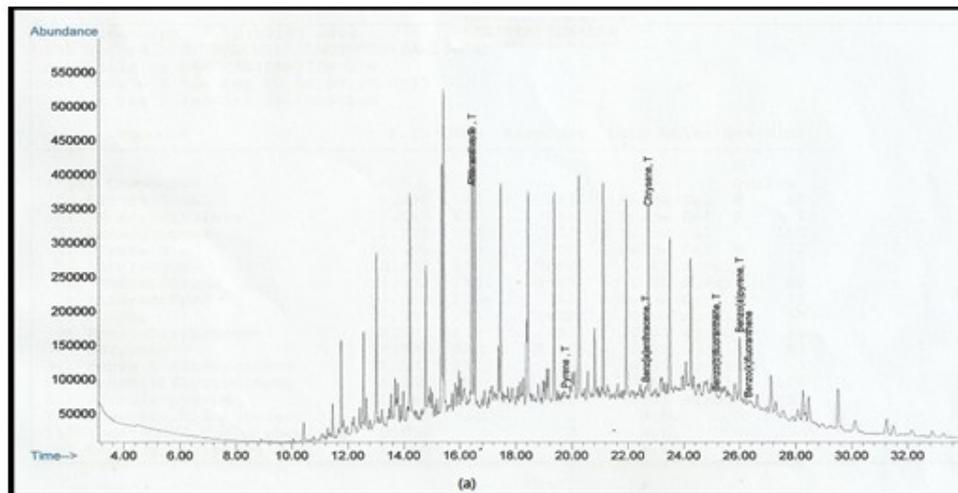
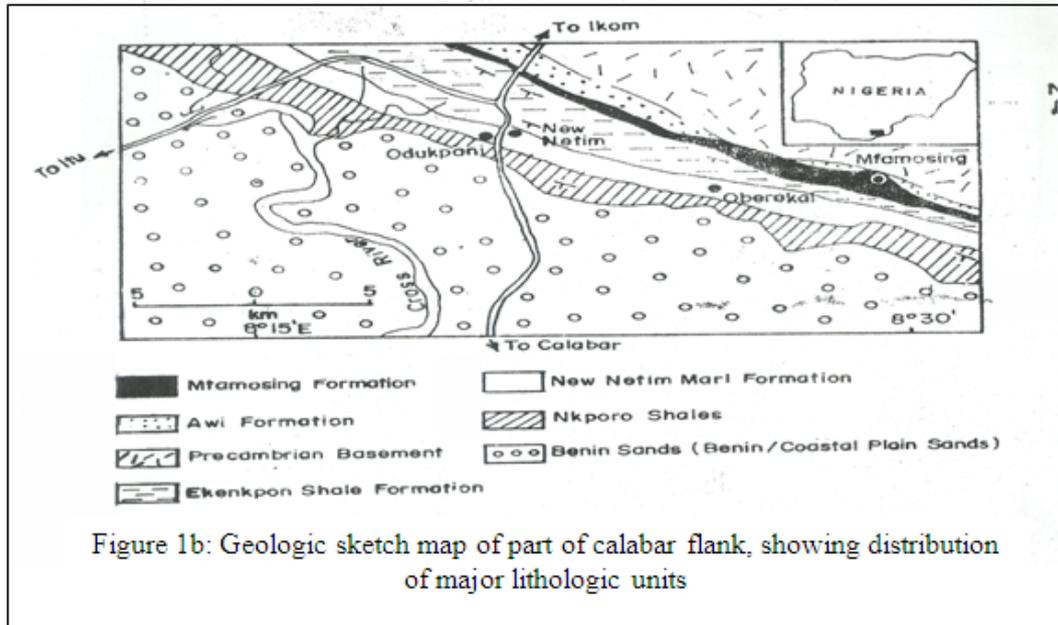
II. Methodology

2.1 Sample Collection

A total of seventeen surface (outcrop) samples of the Cretaceous sediments were picked at various locations where they outcrop (Fig 1b) across the length and breadth of the basin, using appropriate tools and instruments such as hammer and GPS. The outcrop samples were collected and coded as CF-L_nX, where CF is the acronym for the Calabar Flank; L_n, the location serial number and X, to differentiate one sample from another in a particular location. The sediment samples were Awi Sandstone, Mfamosing Limestone, Ekenkpon Shales, New Netim Marl, Nkporo Shale and Nkporo Mudstone.

2.2 Sample Preparation

The outcrop samples were initially cleaned by seeping in 100% dichloromethane with shaking followed by decanting of the solvent until the rocks were clean. After drying, the samples were washed under running tap water and then dried again in an oven at a pre-set temperature of 30° C. About 200mg of the pre-cleaned shale was crushed and accurately weighed into clean LECO crucibles. The rocks were then de-mineralized with hot 10% HCl and afterwards washed repeatedly with distilled water, dried at 60°C, pulverized and preserved.



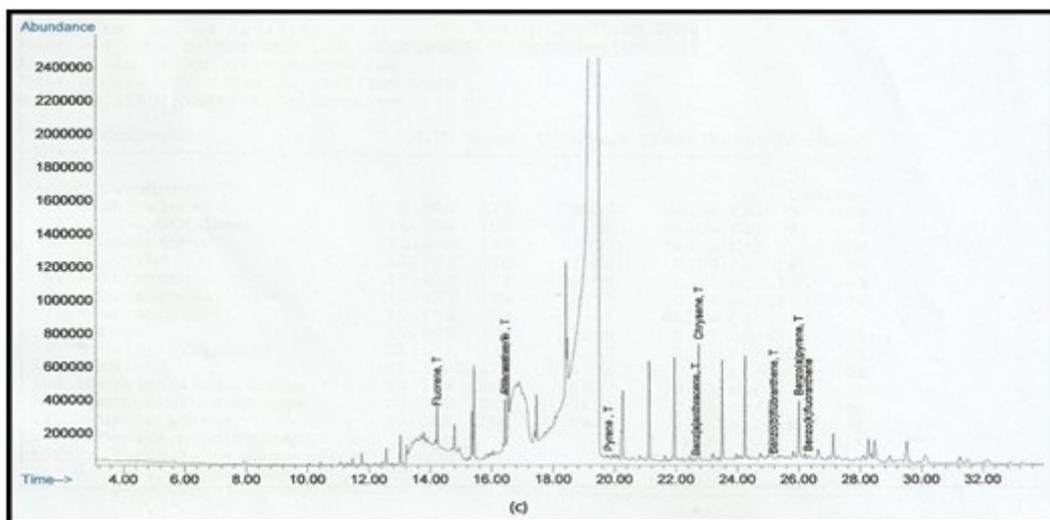


Figure 2c: Representative Mass Chromatogram showing the PAH Profile of Nkporo Shale

Table 1: Lithostratigraphic succession in the Calabar Flank correlated with Anambra Basin (Modified version alter Nyong, 1995)

AGE	CALABAR FLANK BASIN	ANAMBRA BASIN
Maastrichtian	Nkporo Shale	Ikpa, Ajali, Ijebu
Campanian		Ikpa, Ajali, Ijebu
Santonian		
Conagian	New Netim Marl	Agwu Shale
Turonian	Ekenkpon Shale	Eze Ajushale Group
Cenomanian		
Albian	Mfamosing Eormation	Asu River Group
Early Cretaceous	Awi Formation	Basal Grit and Sandstone
Precambrian	Basement Complex of the Oban Massif	

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